

## Inter-calibration of SAR data series for offshore wind resource assessment

Badger, Merete; Ahsbahs, Tobias Torben; Maule, Petr; Karagali, Ioanna

Published in: Remote Sensing of Environment

Link to article, DOI: 10.1016/j.rse.2019.111316

Publication date: 2019

Document Version Peer reviewed version

Link back to DTU Orbit

*Citation (APA):* Badger, M., Ahsbahs, T. T., Maule, P., & Karagali, I. (2019). Inter-calibration of SAR data series for offshore wind resource assessment. *Remote Sensing of Environment*, *232*, Article 111316. https://doi.org/10.1016/j.rse.2019.111316

#### **General rights**

Copyright and moral rights for the publications made accessible in the public portal are retained by the authors and/or other copyright owners and it is a condition of accessing publications that users recognise and abide by the legal requirements associated with these rights.

• Users may download and print one copy of any publication from the public portal for the purpose of private study or research.

- You may not further distribute the material or use it for any profit-making activity or commercial gain
- You may freely distribute the URL identifying the publication in the public portal

If you believe that this document breaches copyright please contact us providing details, and we will remove access to the work immediately and investigate your claim.

| 1<br>2 | Inter-calibration of SAR data series for offshore wind resource assessment   |
|--------|--|
| 3      | Merete Badger <sup>a*</sup> , Tobias Ahsbahs <sup>a</sup> , Petr Maule <sup>a</sup> , Ioanna Karagali <sup>a</sup> |
| 4      |  |
| 5      | <sup>a</sup> Technical University of Denmark, Department of Wind Energy, Frederiksborgvej 399, 4000                |
| 6      | Roskilde, Denmark, e-mail mebc@dtu.dk  |
| 7      | *Corresponding author  |
| 8      |  |
| 9      | Highlights   |
| 10     | • Offshore wind resource assessment requires long-term wind data records.  |
| 11     | • Wind speed retrievals from different European SAR sensors are offset.  |
| 12     | • Biases vary over time and according to scan modes and incidence angles.  |
| 13     | • Inter-calibration can remove biases and improve the accuracy on wind resources.                                  |
| 14     |  |
| 15     | Abstract   |
| 16     | Wind observations in the marine environment are both costly and sparse. This makes wind                            |
| 17     | retrievals from satellite Synthetic Aperture Radar (SAR) an attractive option in connection                        |
| 18     | with planning of offshore wind farms. Because the wind power density is proportional to the                        |
| 19     | wind speed cubed, it is important to achieve the highest possible absolute accuracy on SAR                         |
| 20     | wind speed retrievals for wind energy applications. A method is presented for inter-                               |
| 21     | calibration of SAR observations from Envisat and Sentinel-1A/B. Sensor-specific effects on                         |
| 22     | the SAR-retrieved wind speeds are first quantified through comparisons against collocated                          |
| 23     | ocean buoy observations. Based on global circulation model simulations of wind speed and                           |
| 24     | direction, we retrieve the Normalized Radar Cross Section (NRCS) for different radar                               |
| 25     | incidence angles. Residuals between the retrieved and the observed NRCS are used to inter-                         |
| 26     | calibrate the observed NRCS before reprocessing to SAR wind fields. The inter-calibration                          |

| 27 | leads to an improved agreement between SAR and buoy wind speeds with biases below 0.2 m                  |
|----|--|
| 28 | s <sup>-1</sup> for all investigated SAR sensors. Estimates of the wind resource improve with respect to |
| 29 | the buoy observations for ten of the twelve sites investigated. The average deviation between            |
| 30 | wind power densities is reduced from 20% to 8% as the SAR inter-calibration leads to more                |
| 31 | conservative estimates of the wind resource.   |
| 32 |  |
| 33 | Keywords   |
| 34 | Inter-calibration, offshore wind energy, resource, Synthetic Aperture Radar, Sentinel-1,                 |
| 35 | Envisat  |
|    |  |

37 1. Introduction

38 The Sentinel-1 mission by the European Space Agency (ESA) has secured the availability of Synthetic Apertur Radar (SAR) observations for ocean wind mapping for the years to come. 39 Sentinel-1A (2014-present) and Sentinel-1B (2016-present) are designed for continuation of 40 the previous ESA mission Envisat, which delivered SAR data during 2002-12. SAR 41 instruments are active sensors, which transmit and receive pulses in the microwave range. 42 43 Properties of the ocean surface waves determine the measured return signal. A C-band SAR sensor is sensitive to waves of the cm-scale, which are typically generated by the 44 instantaneous wind stress at the sea surface. 45 46

47 Based on scatterometer observations, empirical relationships have been established between radar backscatter from the sea surface and wind speed at the height 10 m. A similar principle 48 49 can be applied to retrieve wind speeds from SAR observations at a higher spatial resolution and with full coverage over coastal seas (Karagali et al., 2013). Geophysical Model Functions 50 (GMF) for wind speed retrieval at C-band include CMOD4 (Stoffelen and Anderson, 1997), 51 CMOD-IFR2 (Quilfen et al., 1998), CMOD5 (Hersbach et al., 2007), CMOD5.n (Hersbach, 52 2010), CMOD6 and CMOD7 (Stoffelen et al., 2017). The CMOD functions are developed for 53 54 radar observations with vertical polarization in transmit and receive (VV) and a polarization ratio must be applied in order to compensate for the lower signal at HH (Liu et al., 2013; 55 Mouche et al., 2005; Thompson et al., 1998). A new model function called C\_SARMOD2 is 56 57 developed directly from RADARSAT-2 and Sentinel-1 SAR observations (Lu et al., 2018). 58

Wind speed retrievals from Envisat have been compared to in situ observations in different
parts of the world (Chang et al., 2015; Doubrawa et al., 2015; Hasager et al., 2015a; 2015b;
2011; Takeyama et al., 2013a; 2013b) and evaluations of wind speeds from Sentinel-1 are

also published (Ahsbahs et al., 2018; Lu et al., 2018; Monaldo et al., 2016). The Root Mean
Square Error (RMSE) of the SAR wind speed with respect to reference data sets is typically
less than 2.0 m s<sup>-1</sup> whereas the bias can vary largely. The temporal and spatial scales of wind
data should be considered in any comparison analysis (Hasager et al., 2002). Likewise, care
must be taken to compare consistently either the real winds or the Equivalent Neutral Wind
(ENW) (Kara et al., 2008; Portabella and Stoffelen, 2009).

68

The installed wind power capacity is growing rapidly around the world and plans for new 69 installations offshore are ambitious; particularly in Europe and Asia. In order to produce 70 71 robust estimates of the wind resource, the highest possible number of independent wind speed observations is needed. The sampling frequency, which can be achieved from polar-orbiting 72 satellites, is poor compared to the sampling frequencies of typical in situ sensors or numerical 73 74 models. The strength of satellite wind fields lie in the observation of large spatial domains over extensive periods. In order to maximize the number of available satellite wind fields for 75 76 wind resource assessment, the opportunity to combine data series from different sensors is very attractive. However, effects of sensor-specific characteristics need to be taken into 77 account before the data series can be merged. 78

79

Satellite data merging is performed in connection with Climate Data Records (CDRs) defined
as "time series of measurements of sufficient length, consistency and continuity to determine
climate variability and change" (National Research Council, 2004). Merged time-series from
various sensors and for different physical parameters such as ocean surface winds from
scatterometers (Elyouncha and Neyt, 2013; Wentz et al., 2017), ice sheet elevation from
altimeters (Khvorostovsky, 2012), and temperature from microwave sounders (Christy et al.,
1998) already exist. Although the record of wind retrievals from space is not yet long enough

to determine climate variability and change, the community effort is to generate consistent
and stable time-series. Inter-calibration ensures consistency between products from different
sensors and it can be performed using reference data sets of in situ observations and intercomparison among different products (Zeng et al., 2015).

91

92 The objective of this paper is to inter-calibrate SAR observations from Envisat and Sentinel-1 93 SAR and combine them to a single data series suitable for wind speed retrieval and resource assessment. Section 2 describes the data sets analyzed and the pre-processing applied. In 94 Section 3, we present a series of initial comparisons between SAR-retrieved wind speeds and 95 96 ocean buoy observations. A method for inter-calibration of the SAR observations is given in Section 4. In Section 5, comparisons against the reference data set are shown after SAR inter-97 calibration. The effect of SAR inter-calibration on wind resource estimation is examined in 98 99 Section 6. Our findings are discussed in Section 7 and conclusions are given in Section 8.

100

101 2. Data and pre-processing

102

103 2.1 Satellite SAR wind maps

104 This analysis is based on Level-1 SAR data from Envisat and Sentinel-1 A/B, which are available from the Copernicus Open Access Hub at https://scihub.copernicus.eu/. Our focus is 105 on scenes acquired in ScanSAR mode i.e. the Envisat Wide Swath Mode (WSM) and the 106 Sentinel-1 Interferometric Wide Swath (IW) and Extra Wide Swath (EW) Modes. The swath 107 width is fixed at 400 km for WSM and EW and 250 km for IW whereas the length of scenes 108 is variable. All available products covering the seas of Northern Europe (Figure 1) have been 109 downloaded for the period 2002/08/20 to 2018/05/31. Sentinel-1A products generated after 110 2015/11/25 at 10:40 UTC are processed with a radiometric performance enhancement 111

whereas only some of the scenes acquired during the commissioning phase of Sentinel-1A
have been reprocessed (Miranda, 2015). Calibration inconsistencies are therefore still present
for the early Sentinel-1A data. The radiometric accuracy of Sentinel-1B observations has been
satisfactory, and also compatible with that of Sentinel-1A, since launch (Schwerdt et al.,
2017).

117

118 Retrieval of wind speed maps from the Envisat and Sentinel-1 SAR scenes is performed with the SAR Ocean Products System (SAROPS) developed by the Johns Hopkins University, 119 Applied Physics Laboratory and the US National Atmospheric and Oceanographic 120 121 Administration (NOAA) (Monaldo et al., 2014). The CMOD5.n (Hersbach, 2010) function is chosen for the wind speed inversion and the polarization ratio of Mouche et al. (2005) with 122 incidence angle dependence is applied to the scenes acquired in HH. Regardless of the 123 original resolution of satellite SAR products, we average pixels to a size of 0.5 km prior to the 124 wind retrieval processing to reduce effects of random noise and of surface inclination due to 125 longer-period ocean waves. This is common practice for SAR wind retrievals (Dagestad et al., 126 2012). 127

128

129 Because several wind speed and direction pairs may correspond to a single value of backscatter intensity from SAR, information about the wind direction is needed in order to 130 retrieve the wind speed. We obtain the wind directions from the Climate Forecast System 131 Reanalysis (CFSR, http://nomads.ncdc.noaa.gov/data.php?name=access#cfs-reanal-data) 132 during 2002-10 and from the Global Forecast System (GFS) at 0.50° resolution during 2010-133 12 (http://nomads.ncdc.noaa.gov/data/gfsanl) and at 0.25° resolution from 2014 onwards 134 (ftp://ftp.ncep.noaa.gov/pub/data/nccf/com/gfs/prod). The model outputs are interpolated 135 spatially to match the grid cells of the SAR scenes. 136

- 138 Land surfaces are masked out during the SAR wind processing using the Global Self-
- 139 consistent, Hierarchical, High-resolution Geography Database
- 140 (http://www.soest.hawaii.edu/pwessel/gshhg/). Sea ice is detected using the IMS Daily
- 141 Northern Hemisphere Snow and Ice Analysis
- 142 (http://nsidc.org/data/docs/noaa/g02156\_ims\_snow\_ice\_analysis/). The collection of SAR
- 143 wind maps used as the starting point for our analyses is available at
- 144 <u>https://satwinds.windenergy.dtu.dk/</u>.
- 145

146 2.2 Ocean buoy observations

147 Observations from ocean buoys are gathered for the North Sea and part of the North Atlantic

148 for the years 2002 to 2018. To prevent biases, the following criteria are set for buoy stations

to be included in this analysis: *i*) a station must deliver data during the period 2006 to 2017 or

longer; *ii*) no significant change of the buoy position has occurred over time; and *iii*) the buoy

is located at least 10 km from the shoreline. A total of 12 buoy stations live up to the criteria

and these datasets are from three institutions: UK MetOffice (personal communication), the

153 Irish Meteorological Service, Met Éireann

154 (<u>https://erddap.marine.ie/erddap/tabledap/IWBNetwork.html</u>), and the Bundesamt für

155 Seeschifffahrt und Hydrography, BSH (<u>http://nwsportal.bsh.de/</u>).

156

157 The MetOffice and Met Éireann used Ocean Data Acquisition Systems (ODAS) buoys in the

early years and some of them have later been replaced with buoys from the manufacturer

- 159 Fugro. Data from BSH is obtained from light vessels and one moored buoy. Figure 1 shows
- 160 the buoy locations. Position data from the MetOffice buoys are truncated to  $0.1^{\circ}$
- 161 corresponding to an uncertainty of roughly 10 km on the position. The buoy data are quality

- 162 controlled by the respective provider and additional inspection of the time series has been
- 163 performed in connection with this analysis.



Figure 1. The area investigated and positions of the buoys used in this study. The inner
domain shows the area used for SAR inter-calibration in Section 4.

167

The buoy wind speeds and directions are recorded hourly. Measurement heights vary between
3.5 m and 14 m with the vast majority of the observations at heights lower than 10 m. We
extrapolate to 10-m wind speeds using a logarithmic wind profile:

171

172 
$$u(z) = \frac{u_*}{\kappa} \ln \frac{z}{z_0}$$
(1)

173

174 where u(z) is the wind speed at height z (m s<sup>-1</sup>),  $u_*$  is the friction velocity (m s<sup>-1</sup>),  $\kappa$  is the von 175 Kármán constant (~0.4), and  $z_0$  is the surface roughness length, which we set to a constant of 176 0.0002 m.

The air-sea temperature difference, which is needed to estimate atmospheric stability effects, is typically missing in the buoy data sets. We can thus expect a bias on the 10-m wind speed due to the lack of stability correction of the buoy observations. We assume this bias is

181 constant across the Envisat and Sentinel-1 sensing periods.

182

183 3. Initial comparisons of SAR and buoy observations

184 We first compare the wind speeds retrieved from SAR to wind speed observations from the

185 ocean buoys in the North Sea and North Atlantic. The selection criterion for buoy

observations is that their time stamp must be less than 30 minutes from each SAR data

187 acquisition time. To ensure comparability between spatial averaging of the satellite winds and

temporal averaging of the buoy observations, we extract the average SAR wind speeds over

an area of 10 km by 10 km around the buoy positions. We exclude data points where the SAR

190 or buoy wind speeds are below  $0.5 \text{ m s}^{-1}$ .

191

Buoys provide real wind speeds whereas the SAR wind retrievals are expressed as ENW, which are cleaned for atmospheric stability effects and 0.2 m s<sup>-1</sup> higher on average (Kara et al., 2008; Portabella and Stoffelen, 2009). Here, we are primarily interested in the consistency between wind retrievals from Envisat and Sentinel-1. Assuming again that the long-term average stability conditions are similar across sensing periods, we can compare the SAR and buoy wind speeds for this purpose without further correction.

198

Figure 2 shows scatterplots of the buoy wind speeds versus the wind speeds retrieved from SAR. A total of 3099 collocated pairs of Envisat and buoy wind speeds are available and the comparison shows a RMSE of 2.37 m s<sup>-1</sup>. The mean wind speed from Envisat is 0.87 m s<sup>-1</sup>

higher than from the buoy observations. For wind speeds beyond 20 m s<sup>-1</sup>, the SAR wind
speeds are lower than the buoy wind speeds.

| 205 | For the subset of Sentinel-1A scenes acquired during the commissioning phase, the                                     |
|-----|---|
| 206 | comparison show a RMSE of 2.01 m s <sup>-1</sup> and a positive bias for all wind speed bins up to 16 m               |
| 207 | s <sup>-1</sup> . The SAR wind speeds are on average 0.97 m s <sup>-1</sup> higher than the buoy wind speeds. For the |
| 208 | later Sentinel-1A scenes, the RMSE is 1.57 m s <sup>-1</sup> . Comparisons for Sentinel-1B show almost                |
| 209 | similar results with RMSE of 1.58 m s <sup>-1</sup> . For both Sentinel-1 sensors, SAR wind speeds                    |
| 210 | overestimate the buoy wind speeds in the low-wind range. When the wind speed is within the                            |
| 211 | range 7-17 m s <sup>-1</sup> , SAR and buoy wind speeds are almost equal and beyond that, the buoy wind               |
| 212 | speeds are higher. The average bias for Sentinel-1A and B after commissioning is only 0.10-                           |
| 213 | 0.17 m s <sup>-1</sup> and wind speeds from these two sets of SAR observations are very consistent with               |
| 214 | each other. There is an offset with respect to wind retrievals from Envisat and Sentinel-1A                           |
| 215 | observations during commissioning.  |



Figure 2. Comparisons of wind speeds retrieved from SAR against buoy wind speeds for (a)
Envisat; (b) Sentinel-1A commissioning phase; (c) Sentinel-1A; and (d) Sentinel-1B.

219

## 220 3.1 Wind speed dependence on the wind direction input

- 221 To examine the effect of the wind direction input chosen for the SAR wind retrieval
- 222 processing, we repeat the comparisons between SAR and buoy wind speeds using a second
- set of SAR wind speeds retrieved over each of the buoy stations with observed wind

directions from the buoys as input. The buoy wind directions are expected to be more accurate than the model wind directions initially used for the SAR wind retrieval because *i*) they are representative for the exact buoy locations, *ii*) they are measured in a consistent manner across the Envisat and Sentinel-1 sensing periods, and *iii*) they are observed rather than simulated.

229

230 231 Table 1. Summary of comparisons between SAR and buoy wind speeds. The SAR wind speedsare retrieved with wind directions from a model and from buoy observations.

|                              | Envisat |      | Sentinel-1A<br>commissioning<br>phase |      | Sentinel-1A |      | Sentinel-1B |      |
|------------------------------|---------|------|---------------------------------------|------|-------------|------|-------------|------|
| Wind<br>direction<br>input   | Model   | Buoy | Model                                 | Buoy | Model       | Buoy | Model       | Buoy |
| N                            | 3099    | 3099 | 568                                   | 568  | 1660        | 1660 | 1100        | 1100 |
| Bias<br>[m s <sup>-1</sup> ] | 0.87    | 0.92 | 0.94                                  | 0.92 | 0.10        | 0.16 | 0.17        | 0.17 |
| RMSE<br>[m s <sup>-1</sup> ] | 2.37    | 2.37 | 2.01                                  | 1.93 | 1.57        | 1.30 | 1.58        | 1.26 |

Table 1 shows the comparisons of SAR and buoy wind speeds when modelled vs. buoy wind directions are used to drive the wind speed retrieval from SAR. For Envisat, the RMSE is unchanged (2.37 m s<sup>-1</sup>) and the positive bias has increased by 0.05 m s<sup>-1</sup> with respect to the comparison in Figure 2. For all Sentinel-1 data subsets, a small improvement of the RMSE is seen whereas the bias changes by less than 0.1 m s<sup>-1</sup>. The offset between winds from Envisat

and Sentinel-1 after commissioning remains around 0.8 m s<sup>-1</sup> so the quality of wind direction
inputs cannot explain the offsets in wind speed biases between different SAR sensors.

240 Because we find the lowest RMSE for SAR wind speeds retrieved with buoy wind directions,

241 we use these SAR wind retrievals for the remaining part of Section 3.

242

The comparisons presented above all indicate that SAR winds retrieved systematically with 243 244 CMOD5.n overestimate the observed wind speed at low to moderate wind speeds. The positive bias is larger for Envisat and Sentinel-1 during commissioning than for the later 245 Sentinel-1 data series. At high wind speeds, SAR winds retrieved from Envisat and Sentinel-1 246 247 during commissioning still overestimate the observed wind speeds whereas wind speeds 248 retrieved from the later Sentinel-1 data series match the reference wind speeds well. The wind speed biases, which we find for the different SAR sensors and periods, cannot be explained 249 250 by inconsistencies in the ancillary data used to drive the SAR wind retrieval. We therefore turn to examine the effect of different SAR sensing properties on the wind speed accuracy. 251 252

253 3.2 Wind speed dependence on the radar polarization

To investigate the effect of radar polarization on the wind retrieval accuracy, we separate SAR scenes acquired in HH and VV. We can expect the best accuracy at VV polarization since CMOD5.n can be applied directly without a polarization ratio. The majority of SAR scenes in our data set have VV polarization.

258

Table 2 shows results of comparisons between SAR and buoy wind speeds at VV and HH polarization for Envisat and Sentinel-1. The RMSE is significantly lower for VV than HH for all data sets except Sentinel-1B. This is as expected due to the added uncertainty introduced by the polarization ratio we apply to SAR observations acquired with HH-polarization (cf.

Section 1). A positive bias remains for the VV scenes and there is now an average offset of
0.69-0.73 m s<sup>-1</sup> between Envisat and Sentinel-1 retrievals. Envisat scenes acquired in HH
show a large RMSE and a positive bias for all wind speed bins. Sentinel-1 scenes acquired in
HH are associated with a large uncertainty due to the low number of collocated wind speed
samples from SAR and the buoys.

268

Table 2. Summary of comparisons between SAR and buoy wind speeds divided according to
 sensor and polarization.

|                              |      |         | Sentin | el-1A   |      |               |      |             |  |             |  |
|------------------------------|------|---------|--------|---------|------|---------------|------|-------------|--|-------------|--|
|                              | Env  | Envisat |        | Envisat |      | commissioning |      | Sentinel-1A |  | Sentinel-1B |  |
|                              |      |         | phase  |         |      |               |      |             |  |             |  |
| Polarization                 | VV   | HH      | VV     | HH      | VV   | HH            | VV   | HH          |  |             |  |
| N                            | 2777 | 322     | 541    | 61      | 1620 | 42            | 1089 | 11          |  |             |  |
| Bias [m s <sup>-1</sup> ]    | 0.86 | 1.43    | 1.02   | 0.82    | 0.13 | -1.3          | 0.17 | 0.04        |  |             |  |
| RMSE [m<br>s <sup>-1</sup> ] | 2.20 | 3.56    | 2.00   | 3.31    | 1.56 | 2.7           | 1.58 | 1.44        |  |             |  |
|                              |      |         |        |         |      |               |      |             |  |             |  |

271

272 3.3 Wind speed dependence on the radar incidence angle

273 Based on the collocated SAR and buoy wind speed pairs analyzed above, we investigate the

dependence on the SAR-buoy wind speed residuals on the radar incidence angle. Visual

inspection of the SAR derived wind fields indicate that wind speeds can vary across the radar

swath even though the radar incidence angle is taken into account during the SAR wind

277 retrieval. Higher wind speeds typically occur at high incidence angles.

Figure 3 shows the SAR-buoy wind speed residuals as a function of the radar incidence angle. For Envisat, the average wind speed residuals are lower than 1 m s<sup>-1</sup> for incidence angles within the range 20-35°. Below and above this interval, we see a change of the wind speed residuals as a function of incidence angle. The residuals are always positive indicating higher SAR wind speeds compared to the buoy wind speeds. The standard deviation, represented by the error bars, is very high for incidence angles lower than 20°. At all other incidence angles, the standard deviation is  $\pm -2$  m s<sup>-1</sup> or less.

286

Most Sentinel-1 samples are obtained within the incidence angle range of  $30-45^{\circ}$  but a few data points lie within the range of  $20-30^{\circ}$ . During the commissioning phase of Sentinel-1A, we see large fluctuations of the wind speed residuals and error bars of up to +/- 3 m s<sup>-1</sup>. Average wind speed residuals for the later Sentinel-1A acquisitions and for Sentinel-1B are always within the range +/- 1 m s<sup>-1</sup> and the standard deviation remains within +/- 2 m s<sup>-1</sup>. A trend of slightly increasing wind speed residuals with increasing incidence angles is seen in Figure 3 c) and d).

294

Our analyses so far have indicated a consistent difference between wind speed retrievals from Envisat vs. Sentinel-1 A/B, which persists regardless of the wind direction input and the SAR polarization and increases with the SAR incidence angle. To investigate the incidence angle dependence further, we extend the analyses to the Normalized Radar Cross Section (NRCS) input to the SAR wind retrievals.





Figure 3. Residuals between SAR and buoy wind speeds as a function of radar incidence
angle for (a) Envisat; (b) Sentinel-1A commissioning phase; (c) Sentinel-1A; and (d) Sentinel-

1B.

303

# 304 3.4 NRCS dependence on the radar incidence angle

In the following, we use buoy wind speeds and directions to retrieve the NRCS for different radar incidence angles. To achieve this, we apply CMOD5.n in forward mode i.e. we use the buoy wind speed and direction and the radar incidence angle as input and retrieve the NRCS. We then compare the observed and retrieved NRCS.

310 Comparisons of observed and retrieved NRCS from buoy winds are shown in Figure 4. For

- Envisat, the residual of NRCS [dB] is very small at low incidence angles and it increases
- gradually for incidence angles larger than  $20^{\circ}$ . The relationship between the incidence angle
- and the NRCS residuals in dB space is almost linear. For the Sentinel-1A commissioning
- 314 phase, a linear relationship between NRCS residuals and the incidence angle is seen across

the interval 32-41° and there are very few data points at lower incidence angles. For Sentinel1 A/B, the incidence angle range is smaller and the observed NRCS is higher than for Envisat.
This leads to smaller residuals with respect to the retrieved NRCS and again, we see a linear
increase of NRCS with the incidence angle. The results in Figure 4 suggest that changes of
NRCS residuals with the radar incidence angle is the source of the wind speed biases reported
above. In the following, we present a method for correction of the sensor-specific incidence
angle dependence.



Figure 4. Residuals between measured and retrieved NRCS using buoy wind speeds and
directions together with radar incidence angles as input to the simulation for (a) Envisat; (b)
Sentinel-1A commissioning phase; (c) Sentinel-1A; and (d) Sentinel-1B.

329 4. SAR inter-calibration

Inspired by inter-calibration of scatterometers in Elyouncha and Neyt (2013), where sensors
are inter-calibrated using CMOD5.N in forward mode using wind speeds and directions from
global circulation models, we calculate sensor-specific corrections of NRCS. These
corrections are applied to the NRCS observed by different SAR sensors in order to achieve an
inter-calibrated SAR data series.

335

The starting point for the inter-calibration is the set of SAR wind fields obtained within the 336 domain shown in Figure 1 with a distance of at least 20 km from the shore. In addition to the 337 338 10-m wind speed, each data file contains the observed NRCS, radar incidence angle, and look direction as well as wind speeds and directions from a global circulation model (cf. Section 339 2.1). Since model wind speeds and directions are available for all SAR acquisition times and 340 341 all locations, it is convenient to use these for the inter-calibration analysis to achieve the largest possible number of data points for the correction of NRCS. All the listed data layers 342 are resampled to 10 km grid cells to make the SAR observations more comparable to the 343 resolution of the model data and to reduce our computational effort. Resolution cells with 344 wind speeds from either model or SAR-derived winds below 2 m s<sup>-1</sup> and above 20 m s<sup>-1</sup> are 345 filtered out. 346

347

348 NRCS is retrieved from the model wind speed and direction and the radar viewing geometry 349 in a similar fashion as in Section 3.4. Residuals with respect to the observed NRCS (in dB 350 space) are then calculated within incidence angle bins of 1° and a linear fit is made based on 351 the median values. We split our data set according to sensor, polarization, and scan modes. 352 Additionally, we take into account that the calibration of a sensor can change over time by 353 calculating NRCS-corrections on a monthly basis. For each month, data from the previous full

year is used. For the first 12 months a given sensor is in operation, model data covering thesame 12 months are used for correction.

Figure 5 shows examples of the fitted linear functions for one year of data from Envisat,

357 Sentinel-1A commissioning phase, Sentinel-1A, and Sentinel-1B. A clear offset is seen for

358 Envisat, which increases with the incidence angle. The NRCS residuals are less pronounced

359 for Sentinel-1A/B.



360

Figure 5. Examples showing linear fits to the NRCS residual per incidence angle based on
one year of data from (a) Envisat (2008-03 to 2009-03); (b) Sentinel-1A commissioning phase
(2014-11 to 2015-11); (c) Sentinel-1A (2017-04 to 2018-04); and (d) Sentinel-1B (2017-04 to
2018-04).

Subtracting the linear fits from the NRCS observations made by Envisat and Sentinel-1corrects the bias and the slope of NRCS in dB space:

368 
$$\sigma_{IC}^{0}(\theta) = \sigma^{0}(\theta) - fit(year, \theta), \qquad (2)$$

369 where  $\sigma^0$  [dB] is the NRCS and  $\theta$  [°] is the radar incidence angle. The subscript '*IC*' denotes

that NRCS is now inter-calibrated between the sensors. Figure 6 illustrates the entire

371 processing chain of the inter-calibration method applied here.



372

Figure 6. Flow chart showing the processing steps of SAR inter-calibration.

374

373

375 5. Comparisons of SAR and buoy observations after inter-calibration

376 The SAR inter-calibration procedure presented above relies solely on global circulation model

377 wind speeds and directions. We can therefore return to the ocean buoy observations of wind

speed and use these as an independent reference data set. In the following, we compare theinter-calibrated NRCS and SAR wind speed retrievals to the buoy observations.

380

Figure 7 shows residuals between measured and retrieved NRCS as a function of the radar incidence angle. The plots are comparable to those in Figure 4; the only difference being that the NRCS measured from SAR is now inter-calibrated. As a result, residuals of NRCS are very close to zero for the entire span of incidence angles. It is remarkable how the large residuals that we found initially for Sentinel-1 during the commissioning phase are now reduced to a level similar to that of the later Sentinel-1 data series.



387

Figure 7. Residuals between measured and retrieved NRCS after inter-calibration of NRCS
for (a) Envisat; (b) Sentinel-1A commissioning phase; (c) Sentinel-1A; and (d) Sentinel-1B.

390

392 5.1 Wind speed retrieval from corrected NRCS

393 Once we have inter-calibrated the NRCS, we apply CMOD5.n in inverse mode to retrieve wind speeds once again. Wind speed residuals with respect to the buoy observations are 394 395 shown as a function of the radar incidence angle in Figure 8. The plots are comparable to plots in Figure 3 made before the SAR inter-calibration. The inter-calibrated SAR 396 observations lead to much smaller wind speed residuals, especially for Envisat, and there is no 397 398 longer a systematic increase of residuals for increasing incidence angles. 399 It is evident from Figure 7 and Figure 8 that our linear correction of the NRCS works best for 400 radar incidence angles above 25°. At very low incidence angles, few or no Sentinel-1 A/B 401

402 samples are available for fitting a linear function between the radar incidence angle and
403 NRCS residuals. The linear relations found for Envisat at very low incidence angles differ

404 from those found at higher incidence angles. To optimize the wind speed accuracy, we

405 recommend eliminating any data obtained with incidence angles lower than 25°. The

406 following results are calculated with this filter in place.



Figure 8. Residuals between SAR and buoy wind speeds as a function of radar incidence
angle after SAR inter-calibration for (a) Envisat; (b) Sentinel-1A commissioning phase; (c)
Sentinel-1A; and (d) Sentinel-1B.

412 5.2 Effects of inter-calibration on the wind speed accuracy

In Table 2, we saw large differences in the accuracy of default wind speed retrievals from

414 SAR observations acquired with VV and HH polarization. The majority of the HH-polarized

415 SAR scenes in our data set were acquired by Envisat. Table 3 shows how the inter-calibration

- 416 has removed any wind speed bias for retrievals based on Envisat observations with both VV
- and HH polarization. RMSE is also reduced for both VV and HH but its absolute value

418 remains higher for scenes acquired with HH polarization.

Table 3. Summary of comparisons between SAR and buoy wind speeds retrieved from Envisat
observations with VV and HH polarization before and after inter-calibration.

| Polarization              |         | VV         | Н       | Н          |  |
|---------------------------|---------|------------|---------|------------|--|
|                           |         | Inter-     |         | Inter-     |  |
|                           | Default | calibrated | Default | calibrated |  |
| N                         | 1978    | 1978       | 216     | 216        |  |
| Bias [m s <sup>-1</sup> ] | 0.87    | 0.07       | 1.42    | 0.07       |  |
| RMSE [m s <sup>-1</sup> ] | 1.80    | 1.44       | 2.77    | 1.92       |  |

#### 422

423 Figure 9 shows scatter plots of the buoy and SAR wind speeds per sensor after inter-

424 calibration of the NRCS. The number of samples given for each plot is a bit lower than in

425 Figure 2, especially for Envisat. This is due to the filtering of low incidence angles, which

426 was applied in connection with the inter-calibration. In contrast to the plots in Figure 2, we

427 now see a consistency between plots for different SAR sensors. All four plots suggest that

428 SAR winds overestimate buoy observations at low wind speeds up to 7-9 m s<sup>-1</sup> and

429 underestimate with respect to the buoy observations for higher wind speeds.



Figure 9. Comparisons of wind speeds retrieved from inter-calibrated SAR observations
against buoy wind speeds for (a) Envisat; (b) Sentinel-1A commissioning phase; (c) Sentinel1A; and (d) Sentinel-1B.

In Table 4, we present an overview of statistics per SAR sensor before and after the SAR
inter-calibration and using the same set of samples. The inter-calibration consistently leads to

437 a lower RMSE and biases that are close to zero for all sensors.

439 *Table 4. Summary of comparisons between SAR and buoy wind speeds before and after inter-*

|                      | Envisat |        | Sentinel-1A   |          | Sentinel-1A |           | Sentinel-1B |           |
|----------------------|---------|--------|---------------|----------|-------------|-----------|-------------|-----------|
|                      |         |        | commissioning |          |             |           |             |           |
|                      |         |        | pha           | phase    |             |           |             |           |
| Processing           | Default | Inter- | Default       | Inter-   | Default     | Inter-    | Default     | Inter-    |
| choice               |         | calibr |               | calibrat |             | calibrate |             | calibrate |
|                      |         | ated   |               | ed       |             | d         |             | d         |
| N                    | 2194    | 2194   | 551           | 551      | 1659        | 1659      | 1099        | 1099      |
| Bias                 | 0.92    | 0.07   | 0.92          | -0.20    | 0.16        | -0.07     | 0.18        | 0.09      |
| [m s <sup>-1</sup> ] |         |        |               |          |             |           |             |           |
| RMSE                 | 1.92    | 1.49   | 1.93          | 1.55     | 1.30        | 1.26      | 1.26        | 1.24      |
| [m s <sup>-1</sup> ] |         |        |               |          |             |           |             |           |

calibration.

441

The effect of SAR inter-calibration on wind speed retrievals over time is illustrated in Figure 442 443 10. The plot shows how there is a drift of the SAR wind speed accuracy with respect to reference measurements at the buoy stations during Envisat's lifetime. Our correction of 444 NRCS leads to a significant reduction of wind speed residuals during the entire Envisat eera. 445 For Sentinel-1A/B, we see large wind speed residuals for the first two years of operation, 446 which include the commissioning phase of the sensors. The SAR inter-calibration efficiently 447 compensates for wind speed biases so the residuals for Sentinel-1A/B are less than +/-0.2 m s<sup>-</sup> 448 449 <sup>1</sup> at any given time. From the beginning of 2016, the residuals between SAR and reference wind speeds are small and the need for NRCS correction is less pronounced. 450





Figure 10. Residuals of the SAR mean wind speed with respect to buoy observations over
time. The grey curve is based on default SAR wind retrievals and the black curve is based on
wind retrievals from inter-calibrated SAR observations.

Our results indicate that we have successfully removed biases on wind retrievals from the different SAR sensors. The bias removal is crucial for merging of the wind speeds retrieved from Envisat and Sentinel-1A/B to a single time series, which is desired for e.g. wind energy resource assessment. In the following, we will examine the effect of inter-calibration on the wind resource we can estimate for each of the buoy locations.

461

462 6. Wind resource assessment

463 The principle of satellite based wind resource mapping is similar to that of wind resource

464 assessment from time series observations e.g. with a meteorological mast (Troen and

Petersen, 1989) or from outputs of numerical models (Hahmann et al., 2015). For a given grid cell, a time series of SAR wind samples can be constructed and analyzed statistically. A Weibull function is fitted to the frequency distribution of wind speed bins. The function is defined by a scale parameter, *A* and a shape parameter, *k*. From these, the wind power density, E (W m<sup>-2</sup>) is calculated:

470

471 
$$E = \frac{1}{2} \rho A^3 \Gamma \left( 1 + \frac{3}{k} \right), \tag{3}$$

472

473 where  $\rho$  is the air density (here set to 1.23 kg m<sup>-3</sup>). Repeating this analysis for each point in a 474 geographical grid will lead to wind resource maps (Badger et al., 2010; Doubrawa et al., 475 2015; Hasager et al., 2015)

476

In order to examine the effect of SAR inter-calibration on wind resource estimates, we 477 calculate the wind power density for each buoy location. The wind power densities are listed 478 in together with the residuals between SAR and buoy wind resources before and after inter-479 calibration of the SAR data sets. For ten of the 12 buoy locations, we find that the wind power 480 density estimated from SAR after inter-calibration shows a lower bias with respect to the 481 482 buoy observations. The average numerical deviation from the buoy observations is 20% before and 8% after SAR inter-calibration. 483 484 485 486 487

|                      | 2   |                |                     |                |
|----------------------|---|----------------|---------------------|----------------|
| Tull 5 Wind a second | 1   | f              | internation and a d | 1              |
| I ADIE V WINA DOWER  | $n\rho n \sin \rho \sin \rho \sin \nu m^{-1}$ | τον της τωρινό | ιηνρετισμτρμ        | nuov iocations |
|                      |   |                | <i>miresulture</i>  | onoy iocunons. |

| Station   | N   | E <sub>buoy</sub> | E <sub>SAR</sub> | E <sub>SAR_IC</sub> | E <sub>SAR</sub> - E <sub>buoy</sub> | E <sub>SAR_IC</sub> - E <sub>buoy</sub> |
|-----------|-----|-------------------|------------------|---------------------|--------------------------------------|---|
| BRITTANY  | 735 | 559               | 632              | 545                 | 73                                   | -14                                     |
| 62091     | 644 | 506               | 582              | 486                 | 76                                   | -20                                     |
| GASCOIGNE | 557 | 450               | 466              | 399                 | 16                                   | -51                                     |
| K7        | 496 | 825               | 948              | 784                 | 123                                  | -41                                     |
| TWEms     | 475 | 515               | 595              | 501                 | 80                                   | -14                                     |
| 62093     | 456 | 638               | 839              | 712                 | 201                                  | 74                                      |
| 62094     | 449 | 500               | 611              | 489                 | 111                                  | -11                                     |
| DtBucht   | 441 | 461               | 565              | 442                 | 104                                  | -19                                     |
| NsbII     | 383 | 681               | 598              | 523                 | -83                                  | -158                                    |
| 62092     | 276 | 514               | 719              | 563                 | 205                                  | 49                                      |
| K1        | 260 | 778               | 895              | 727                 | 117                                  | -51                                     |
| K5        | 222 | 819               | 1013             | 885                 | 194                                  | 66                                      |
| K2        | 109 | 770               | 994              | 872                 | 224                                  | 102                                     |

It is not clear why the inter-calibration leads to higher residuals at the two sites Gascoigne and
Nsbll. One explanation could be that the fitting of a Weibull function introduces some
uncertainty to the wind resource estimation. In fact, when we calculate a simple mean value of
the wind speed observations, the two stations show better agreement with the reference data
after inter-calibration. Other possible reasons for the deviation at the two stations could be
issues with the buoy data quality e.g. inaccurate positioning, instrument faults, or biases
caused by the vertical extrapolation of wind speed observations.



Figure 11. Weibull distributions for the two sites Brittany (a-b) and 62091 (c-d). The Weibull
distributions are shown before and after SAR inter-calibration.



parameter is also much closer to the buoy observations after inter-calibration and the absolute
residual of the wind power density improves from 73 to 14 W m<sup>-2</sup>.

514

At the buoy station 62091, prevailing winds are from more southerly directions due to channeling effects within the Irish Sea. The difference between Weibull curves before and after inter-calibration of the SAR data is less pronounced than for Brittany. In fact, the values of Weibull *k* are identical to the buoy observations before inter-calibration whereas a difference of 0.08 is found after inter-calibration. As for Brittany, we find that Weibull *A* and the wind power density is reduced significantly after the SAR inter-calibration.

521

Table 6. Summary of the bias, RMSE, and MAE of wind resource assessments averaged for
the 12 buoy stations investigated. The mean wind speed (U), wind power density (E), Weibull
scale (A), and shape (k) parameters are calculated before and after the inter-calibration of
SAR observations.

|      | $U [{ m m s}^{-1}]$ |            | <i>E</i> [W m <sup>-2</sup> ] |            | <i>A</i> [m s <sup>-1</sup> ] |            | k [-]   |            |
|------|---------------------|------------|-------------------------------|------------|-------------------------------|------------|---------|------------|
|      | Default             | Inter-     | Default                       | Inter-     | Default Inter-                |            | Default | Inter-     |
|      |                     | calibrated |                               | calibrated |                               | calibrated |         | calibrated |
| Bias | 0.60                | 0.05       | 111                           | -7         | 0.49                          | -0.10      | -0.05   | -0.06      |
| RMSE | 0.67                | 0.27       | 138                           | 65         | 0.57                          | 0.28       | 0.16    | 0.12       |
| MAE  | 0.61                | 0.21       | 124                           | 52         | 0.50                          | 0.26       | 0.13    | 0.11       |

526

The bias, RMSE, and Mean Absolute Error (MAE) averaged for all 12 buoy stations are
summarized in Table 6. The bias on *U* is reduced to almost zero and this reduces the bias on
both *E* and Weibull-*A* significantly. All three biases change from positive to negative values

after the SAR inter-calibration and this leads to more conservative estimates of the wind resource. The bias on Weibull-*k* remains the same. The RMSE is also reduced for U, E, and Weibull-*A* indicating a lower uncertainty of wind resource estimates after the SAR intercalibration.

535

536 7. Discussion

537 Our initial processing of wind speed maps from Envisat and Sentinel-1A/B observations lead to a positive bias for all the SAR sensors investigated but with a large offset between Envisat 538 and Sentinel-1A/B. This is critical if a long time series based on all available SAR 539 540 observations is desired e.g. for wind resource assessment. The RMSE found in our initial comparisons with buoy observations of wind speed are similar to values found in previous 541 studies based on Envisat (Chang et al., 2015; Doubrawa et al., 2015; Hasager et al., 2015a; 542 543 2015b; 2011; Takeyama et al., 2013a; 2013b) and Sentinel-1A/B (Ahsbahs et al., 2018; Lu et al., 2018; Monaldo et al., 2016). Our analyses confirm that observations from the two 544 Sentinel-1 sensors A and B lead to wind speeds having almost the same level of accuracy with 545 respect to reference data sets if the commissioning phase of the Sentinel-1A data series is 546 neglected. 547

548

Our analyses show for the first time how observations from different SAR sensors can be inter-calibrated in the same fashion as scatterometer observations are inter-calibrated in connection with CDR development (cf. Section 1). So far, efforts to inter-calibrate SAR observations from different sensors have been limited since relatively few users of the observations see a need for long-term climatological variables. Efforts have instead been dedicated to determining the most suitable GMF for SAR wind retrieval in different areas of the world (Christiansen et al., 2006; Hasager et al., 2015; Takeyama et al., 2013b). Our results

indicate that a single GMF cannot retrieve wind speeds from multiple sensors accurately as
long as NRCS residuals vary according to sensor type, scan mode, incidence angle, and over
the sensor lifetime. It is thus necessary to inter-calibrate the NRCS before wind retrieval
processing unless a new GMF is developed specifically for the SAR sensors in question so
that inter-calibration is indirectly performed through tuning of the GMF (Lu et al., 2018).

561

562 The inter-calibration method presented here leads to a significant reduction of the offset between wind speed retrievals from Envisat and Sentinel-1A/B observations. After inter-563 calibration, the average wind speed bias does not exceed  $\pm -0.20$  m s<sup>-1</sup> for any sensor 564 investigated here and the RMSE on wind speeds is less than 1.55 m s<sup>-1</sup> with respect to ocean 565 buoy observations. For Sentinel-1A/B, we achieve almost zero wind speed bias and a RMSE 566 as low as 1.24 m s<sup>-1</sup>. The difference between wind resource estimates from SAR and the buoy 567 568 wind speeds is reduced as a result of inter-calibration for ten of the 12 sites investigated. The inter-calibration removes positive biases from the SAR observations and this leads to lower 569 570 and more conservative estimates of the wind power density. From an industry perspective, it is important to operate with conservative rather than over-optimistic resource estimates to 571 ensure that potential new wind farms can deliver on feasibility as expected. 572

573

This work relies on several assumptions, which may be investigated further in future research. Wind speed retrievals using CMOD5.n result in the ENW, which is offset from the real wind speed (Kara et al., 2008; Portabella and Stoffelen, 2009). Over the seas of Northern Europe, this offset is found to be smaller than 0.1 m s<sup>-1</sup> for the height 10 m and it increases for higher levels in the atmosphere (Badger et al., 2016). Our comparisons between SAR and model wind speeds and the calculation of NRCS corrections do not take the offset between ENW and real winds into account. We assume the offset to be constant over time from the Envisat

to the Sentinel-1A/B era and so, the impact will be constant for all the SAR data sets
investigated. In reality, the atmospheric stability has a seasonal variation as it is temperaturedriven. A seasonal inter-calibration analysis would be helpful for quantifying the effect of

584 atmospheric stability.

585

586 In connection with the fitting of linear functions to calculate NRCS corrections, we also 587 assume that the modelled wind speeds will on average converge to the true mean wind speed (both spatially and temporally); otherwise we are adjusting to an offset wind speed. 588 Comparisons between model and in situ wind speeds (not shown here) indicate that the model 589 590 simulations are indeed consistent with the real wind speeds in the long-term. Our linear fitting is performed for the wind speed interval 2-20 m s<sup>-1</sup>. A high uncertainty is anticipated for 591 extremely low and high wind speeds due to lower sampling rates and a saturation problem of 592 593 GMFs at high wind speeds. Work is ongoing in the satellite wind community to resolve extremely high wind speeds thanks to the availability of new cross-polarized SAR sensors 594 595 (Mouche et al., 2017; Zadelhoff et al., 2014). Further developments of our inter-calibration method might take high wind speeds better into account. 596

597

598 The spatial and temporal collocation of data sets in our analyses add uncertainties to our 599 findings because: i) model simulations and buoy observations are available every hour and the offset in time from the SAR observations may thus be up to 30 minutes; *ii*) the exact geo-600 location of ocean buoys can be difficult to determine from the metadata provided with the 601 602 wind speed data; and *iii*) the measurement height for the buoy winds may not be accurate and interpolation to the height of 10 m adds additional uncertainty to wind speed estimates. In 603 604 order to examine the robustness of our inter-calibration method, it would be valuable to test it for other independent sites where high-quality wind observations are available. The ideal test 605

site would provide offshore wind measurements at the height 10 m together with air-seatemperature differences suitable for atmospheric stability correction.

608

609 The successful inter-calibration of SAR data from the European Space Agency presented here could potentially be extended to cover other SAR sensors and scan modes. As an example, 610 611 long C-band SAR data series have been acquired by Radarsat-1/2, which is soon to be 612 continued with the Radarsat Constellation Mission. Sensors operating at X-band or L-band represent other possible extensions of the data series investigated here. An added benefit of 613 using SAR observations from a variety of sensors in combination would be that diurnal wind 614 615 speed variability can be better resolved. 616 617 At present, the calibration of individual SAR sensors is the responsibility of different space 618 agencies and it is typically governed by different requirements. The method for intercalibration described here can be applied by any end user of SAR data and it is thus promising 619 620 for inter-calibration of multiple SAR data sets obtained in the past, present and future. Potentially, an inter-calibrated long-term record of SAR wind speeds could be established and 621 622 offered through publicly available data portals. This would facilitate the best possible 623 accuracy on long-term average wind speeds offshore for many applications including wind 624 energy resource assessment.

625

626 8. Conclusion

We have presented a method for inter-calibration of SAR observations with the purpose of constructing a long-term record of wind speed retrievals from SAR. Correction of the NRCS prior to wind retrieval processing efficiently removes biases on wind speeds from Envisat and Sentinel-1A/B observations. The correction varies according to the SAR sensor, scan mode,

| 631 | radar incidence angle, and also over the sensor lifetime. The inter-calibration leads to a |
|-----|--|
| 632 | significant reduction of wind speed biases and uncertainties expressed through the RMSE.   |
| 633 | Wind resource estimates become more conservative as a result of the SAR inter-calibration. |
| 634 | Our successful calculation of a long-term wind speed record form SAR observations is       |
| 635 | promising and has a potential for extension using other SAR sensors from the past, present |
| 636 | and future. Ultimately, this could lead to establishment of a new derived product offering |
| 637 | long-term SAR wind data for wind energy resource assessment and other applications.        |
| 638 |  |
| 639 | Acknowledgements   |
| 640 | This work received funding from the EU H2020 program under grant agreement no. 730030      |
| 641 | (CEASELESS project).   |
| 642 |  |
| 643 | References   |
| 644 | Ahsbahs, T., Badger, M., Karagali, I., Larsén, X.G., 2017. Validation of Sentinel-1A SAR   |
| 645 | coastal wind speeds against scanning LiDAR. Remote Sens. 9.                                |
| 646 | https://doi.org/10.3390/rs9060552  |
| 647 | Ahsbahs, T., Badger, M., Volker, P., Hansen, K.S., Hasager, C.B., 2018. Applications of    |
| 648 | satellite winds for the offshore wind farm site Anholt. Wind Energy Sci. 3, 573–588.       |
| 649 | https://doi.org/10.5194/wes-2018-2   |
| 650 | Badger, M., Badger, J., Nielsen, M., Hasager, C.B., Peña, A., 2010. Wind Class Sampling of |
| 651 | Satellite SAR Imagery for Offshore Wind Resource Mapping. J. Appl. Meteorol.               |
| 652 | Climatol. https://doi.org/10.1175/2010JAMC2523.1   |
| 653 | Badger, M., Peña, A., Hahmann, A.N., Mouche, A.A., Hasager, C.B., 2016. Extrapolating      |
| 654 | satellite winds to turbine operating heights. J. Appl. Meteorol. Climatol. 55.             |
| 655 | https://doi.org/10.1175/JAMC-D-15-0197.1   |

- 656 Chang, R., Zhu, R., Badger, M., Hasager, C., Xing, X., Jiang, Y., 2015. Offshore Wind
- 657 Resources Assessment from Multiple Satellite Data and WRF Modeling over South

- 659 Christiansen, M.B., Koch, W., Horstmann, J., Hasager, C.B., Nielsen, M., 2006. Wind
- resource assessment from C-band SAR. Remote Sens. Environ. 105, 68–81.
- 661 Christy, J.R., Spencer, R.W., Lobl, E.S., 1998. Analysis of the merging procedure for the
- 662 MSU daily temperature time series. J. Clim. 11, 2016–2041.
- 663 https://doi.org/10.1175/1520-0442-11.8.2016
- 664 Dagestad, K.-F., Horstmann, J., Mouche, A., Perrie, W., Shen, H., Zhang, B., Li, X.,
- 665 Monaldo, F., Pichel, W., Lehner, S., Badger, M., Hasager, C.B., Furevik, B., Foster,
- R.C., Falchetti, S., Caruso, M., Vachon, P., 2012. Wind retrieval from synthetic aperture
- radar an overview, in: Proceedings of SEASAR 2012 Advances in SAR Oceanography,
  Tromsø Norway.
- 669 Doubrawa, P., Barthelmie, R.J., Pryor, S.C., Hasager, C.B., Badger, M., Karagali, I., 2015.
- 670 Satellite winds as a tool for offshore wind resource assessment: The Great Lakes Wind
- 671 Atlas. Remote Sens. Environ. 168. https://doi.org/10.1016/j.rse.2015.07.008
- Elyouncha, A., Neyt, X., 2013. C-band satellite scatterometer intercalibration. IEEE Trans.
- 673 Geosci. Remote Sens. 51, 1478–1491. https://doi.org/10.1109/TGRS.2012.2217381
- Hahmann, A.N., Vincent, C.L., Peña, A., Lange, J., Hasager, C.B., 2015. Wind climate
- estimation using WRF model output: Method and model sensitivities over the sea. Int. J.
- 676 Climatol. 35, 3422–3439. https://doi.org/10.1002/joc.4217
- Hasager, C.B., Badger, M., Nawri, N., Furevik, B.R., Petersen, G.N., Bjornsson, H., Clausen,
- 678 N.-E., 2015a. Mapping Offshore Winds Around Iceland Using Satellite Synthetic
- Aperture Radar and Mesoscale Model Simulations. IEEE J. Sel. Top. Appl. Earth Obs.
- 680 Remote Sens. 8. https://doi.org/10.1109/JSTARS.2015.2443981

<sup>658</sup> China Sea. Remote Sens. 7, 467–487. https://doi.org/10.3390/rs70100467

- Hasager, C.B., Badger, M., Peña, A., Larsén, X.G., Bingöl, F., 2011. SAR-based wind
- resource statistics in the Baltic Sea. Remote Sens. 3. https://doi.org/10.3390/rs3010117
- Hasager, C.B., Dellwik, E., Nielsen, M., Furevik, B., 2002. Validation of ERS-2 SAR

684 offshore wind-speed maps in the North Sea. Int. J. Remote Sens.

- Hasager, C.B., Mouche, A., Badger, M., Bingöl, F., Karagali, I., Driesenaar, T., Stoffelen, A.,
- 686 Peña, A., Longépé, N., 2015b. Offshore wind climatology based on synergetic use of
- 687 Envisat ASAR, ASCAT and QuikSCAT. Remote Sens. Environ. 156.
- 688 https://doi.org/10.1016/j.rse.2014.09.030
- Hasager, C.B., Mouche, A., Badger, M., Bingöl, F., Karagali, I., Driesenaar, T., Stoffelen, A.,
- 690 Peña, A., Longépé, N., 2015. Offshore wind climatology based on synergetic use of
- 691 Envisat ASAR, ASCAT and QuikSCAT. Remote Sens. Environ. 156, 247–263.
- 692 https://doi.org/10.1016/j.rse.2014.09.030
- 693 Hersbach, H., 2010. Comparison of C-Band Scatterometer CMOD5.N Equivalent Neutral
- 694 Winds with ECMWF. J. Atmos. Ocean. Technol. 27, 721–736.
- 695 https://doi.org/10.1175/2009JTECHO698.1
- Hersbach, H., Stoffelen, A., de Haan, S., 2007. An improved C-band scatterometer ocean
  geophysical model function: CMOD5. J. Geophys. Res. 112, 16 pp.
- Kara, A.B., Wallcraft, A.J., Bourassa, M.A., 2008. Air-sea stability effects on the 10 m winds
- 699 over the global ocean: Evaluations of air-sea flux algorithms. J. Geophys. Res. Ocean.
- 700 113, 1–14. https://doi.org/10.1029/2007JC004324
- Karagali, I., Larsén, X., Badger, M., Peña, A., Hasager, C., 2013. Spectral Properties of
- 702 ENVISAT ASAR and QuikSCAT Surface Winds in the North Sea. Remote Sens. 5,
- 703 6096–6115. https://doi.org/10.3390/rs5116096
- Khvorostovsky, K.S., 2012. Merging and analysis of elevation time series over greenland ice
- sheet from satellite radar altimetry. IEEE Trans. Geosci. Remote Sens. 50, 23–36.

- 706 https://doi.org/10.1109/TGRS.2011.2160071
- Liu, G., Yang, X., Li, X., Zhang, B., Pichel, W., Li, Z., Zhou, X., 2013. A Systematic
- 708 Comparison of the Effect of Polarization Ratio Models on Sea Surface Wind Retrieval
- From C-Band Synthetic Aperture Radar. IEEE J. Sel. Top. Appl. Earth Obs. Remote
- 710 Sens. 6, 1100–1108. https://doi.org/10.1109/JSTARS.2013.2242848
- Lu, Y., Zhang, B., Perrie, W., Mouche, A.A., Li, X., Wang, H., 2018. A C-Band geophysical
- model function for determining coastal wind speed using synthetic aperture radar. IEEE
- 713 J. Sel. Top. Appl. Earth Obs. Remote Sens. 11, 2417–2428.
- 714 https://doi.org/10.1109/JSTARS.2018.2836661
- 715 Miranda, N., 2015. S-1A TOPS Radiometric Calibration Refinement # 1.
- 716 Monaldo, F., Jackson, C., Li, X., Pichel, W.G., 2016. Preliminary Evaluation of Sentinel-1A
- Wind Speed Retrievals. IEEE J. Sel. Top. Appl. Earth Obs. Remote Sens. 9, 2638–2642.
  https://doi.org/10.1109/JSTARS.2015.2504324
- 719 Monaldo, F.M., Li, X., Pichel, W.G., Jackson, C.R., 2014. Ocean Wind Speed Climatology
- from Spaceborne SAR Imagery. Bull. Am. Meteorol. Soc. 95, 565–569.
- 721 https://doi.org/10.1175/BAMS-D-12-00165.1
- Mouche, A.A., Chapron, B., Zhang, B., Husson, R., 2017. Combined Co- and Cross-Polarized
- 723 SAR Measurements Under Extreme Wind Conditions. IEEE Trans. Geosci. Remote

724 Sens. 55, 6746–6755. https://doi.org/10.1109/TGRS.2017.2732508

- Mouche, A.A., Hauser, D., Daloze, J.F., Guerin, C., 2005. Dual-polarization measurements at
- 726 C-band over the ocean: Results from airborne radar observations and comparison with
- ENVISAT ASAR data. IEEE Trans. Geosci. Remote Sens. 43, 753–769.
- 728 National Research Council, 2004. Climate Data Records from Environmental Satellites:
- 729 Interim Report. The National Academies Press, Washington DC.
- 730 https://doi.org/10.17226/10944

- Portabella, M., Stoffelen, a., 2009. On Scatterometer Ocean Stress. J. Atmos. Ocean.
  Technol. 26, 368–382. https://doi.org/10.1175/2008JTECHO578.1
- 733 Quilfen, Y., Chapron, B., Elfouhaily, T., Katsaros, K., Tournadre, J., 1998. Observation of
- tropical cyclones by high-resolution scatterometry. J. Geophys. Res. 103, 7767–7786.
- 735 Schwerdt, M., Schmidt, K., Ramon, N.T., Klenk, P., Yague-Martinez, N., Prats-Iraola, P.,
- Zink, M., Geudtner, D., 2017. Independent system calibration of sentinel-1B. Remote

737 Sens. 9, 1–34. https://doi.org/10.3390/rs9060511

738 Stoffelen, A., Anderson, D.L.T., 1997. Scatterometer data interpretation: Estimation and

validation of the transfer function CMOD4. J. Geophys. Res. 102, 5767–5780.

- 740 Stoffelen, A., Verspeek, J.A., Vogelzang, J., Verhoef, A., 2017. The CMOD7 Geophysical
- 741 Model Function for ASCAT and ERS Wind Retrievals. IEEE J. Sel. Top. Appl. Earth
- 742 Obs. Remote Sens. 10, 2123–2134. https://doi.org/10.1109/JSTARS.2017.2681806
- Takeyama, Y., Ohsawa, T., Kozai, K., Hasager, C.B., Badger, M., 2013. Comparison of
- geophysical model functions for SAR wind speed retrieval in japanese coastal waters.
  Remote Sens. 5. https://doi.org/10.3390/rs5041956
- Takeyama, Y., Ohsawa, T., Kozai, K., Hasager, C.B., Badger, M., 2013. Effectiveness of
- 747 WRF wind direction for retrieving coastal sea surface wind from synthetic aperture
  748 radar. Wind Energy 16, 865–878.
- 749 Takeyama, Y., Ohsawa, T., Kozai, K., Hasager, C.B., Badger, M., 2010. Effect of Wind
- 750 Direction on ENVISAT ASAR Wind Speed Retrieval, in: Proceedings (CD-ROM).
- 751 Techno-Ocean Network, p. 8.
- 752 Thompson, D., Elfouhaily, T., Chapron, B., 1998. Polarization ratio for microwave
- backscattering from the ocean surface at low to moderate incidence angles. pp. 1671–
  1676.
- 755 Troen, I., Petersen, E.L., 1989. European Wind Atlas. Risø National Laboratory, Roskilde,

756 Denmark.

| 757 | Zadelhoff, G.J. Van, Stoffelen, A., Vachon, P.W., Wolfe, J., Horstmann, J., Belmonte Rivas,     |
|-----|---|
| 758 | M., 2014. Retrieving hurricane wind speeds using cross-polarization C-band                      |
| 759 | measurements. Atmos. Meas. Tech. 7, 437-449. https://doi.org/10.5194/amt-7-437-2014             |
| 760 | Zeng, Y., Su, Z., Calvet, J.C., Manninen, T., Swinnen, E., Schulz, J., Roebeling, R., Poli, P., |
| 761 | Tan, D., Riihelä, A., Tanis, C.M., Arslan, A.N., Obregon, A., Kaiser-Weiss, A., John,           |
| 762 | V.O., Timmermans, W., Timmermans, J., Kaspar, F., Gregow, H., Barbu, A.L.,                      |
| 763 | Fairbairn, D., Gelati, E., Meurey, C., 2015. Analysis of current validation practices in        |
| 764 | Europe for space-based climate data records of essential climate variables. Int. J. Appl.       |
| 765 | Earth Obs. Geoinf. 42, 150–161. https://doi.org/10.1016/j.jag.2015.06.006                       |
|     |   |